# *P*-wave velocity test for assessment of geotechnical properties of some rock materials

# SAFFET YAGIZ

Faculty of Engineering, Department of Geological Engineering, Pamukkale University, Denizli, Turkey

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Abstract. *P*-wave velocity test, a non-destructive and easy method to apply in both field and laboratory conditions, has increasingly been conducted to determine the geotechnical properties of rock materials. The aim of this study is to predict the rock properties including the uniaxial compressive strength, Schmidt hardness, modulus of elasticity, water absorption and effective porosity, slake durability index, saturated and dry density of rock using *P*-wave velocity ( $V_p$ ). For this purpose geotechnical properties of nine different rock types were determined in the laboratory and their mineralogical composition examined using thin section analysis. Utilizing the generated data, sets of empirical equations were developed between  $V_p$  and relevant quantified rock parameters. The validity of the obtained empirical equations was confirmed using statistical analysis. It is evident that rock texture and mineralogical compositions affect the geotechnical properties of rock materials. Therefore, the best relationship obtained between both E and UCS with  $V_p$  in the correlation coefficient of 0.92 and 0.95 in that order. It is concluded that  $V_p$  could be practically used for estimating the measured rock properties except dry and saturated density of rocks (r = 0.58 and 0.46 respectively).

Keywords. Empirical equations; P-wave velocity; regression analysis; rock properties.

## 1. Introduction

*P*-wave velocity test that can be carried out both in the laboratory and on-site is a common non-destructive testing method used in civil, geotechnical and mining projects such as underground opening, quarrying, blasting and ripping. Seismic techniques can be used for predicting the rock mass deformation and stress as well as extend of damage zone developed around the underground opening and tunnels (Onodera 1963; Hudson et al 1980; Gladwin 1982). The method is also commonly used for determination of rock weathering degree and rock mass characterization (Turk and Dearman 1986; Karpuz and Pasamehmetoglu 1997; Boadu 1997). Thill and Bur (1969) stated that the P-wave velocity changes with porosity and degree of saturation. Lama and Vutukuri (1978) indicated that the wetting of rock usually leads to a rise in the *P*-wave velocity. Several researchers (Hawkins and McConnell 1992; Ulusay et al 1994; Tugrul and Zarif 1999; Kahraman 2001; Yasar and Erdogan 2004; Kahraman and Yeken 2008; Sharma and Singh 2007; Yagiz 2009) reported that the  $V_{\rm p}$  has relationship with some rock properties such as uniaxial compressive strength, hardness, density and slake durability index of rock as shown in table 1. However, obtained correlations are not constant and can be varied with rock types. This paper attempts to investigate the empirical relationship between  $V_p$  and rock properties including the uniaxial compressive strength (UCS), modulus of elasticity (E), Schmidt hardness (Hr), slake durability index (Id<sub>2</sub>), effective porosity (n'), water absorption by weight (w), and both saturated ( $\rho_{sat}$ ) and dry ( $\rho_{dry}$ ) density of rocks. Further, obtained results are also compared with previous studies in the literature.

#### 2. Rock sampling and laboratory tests

Rock blocks were collected from various stone quarries located around the cities of Denizli and Antalya in southwestern Turkey (figure 1). Nine different rocks consisting of four types of travertine, three types of limestone and two types of schist were collected from the study area. Each rock block was inspected to ensure that it would provide standard testing specimens without macroscopic defects, alteration zones and fractures. Following the European Norms (EN 2000a, b) and ISRM (1981) suggested methods, relevant rock properties i.e.  $V_{\rm p}$ , UCS, E, Hr, Id<sub>2</sub>, n', w,  $\rho_{\rm sat}$  and  $\rho_{\rm dry}$  were determined. For each test, 10 rock samples were prepared. The average values obtained along with the standard deviation are given in table 2. Entire tests were performed on intact rock samples. If a rock failed along the anisotropy zone or weak-

<sup>(</sup>syagiz@pau.edu.tr)

Researchers	Equations	<i>r</i> value	Rock type/ lithology	UCS (MPa) $\rho(g/cm^3)$	$V_{\rm p}({\rm km/s})$
Tugrul and Zarif (1999) Kahraman (2001) Yasar and Erdogan (2004)	UCS = $35 \cdot 54 \cdot V_p - 55$ UCS = $9 \cdot 95 \cdot V_p^{1.21}$ UCS = $(V_p - 2 \cdot 0195)/0.032$ $\rho = (V_p + 7 \cdot 707)/4 \cdot 3183$	0.80 0.83 0.81 0.80	Igneous rocks Limestone, marble Lime, marble, dolomite	100–200 10–160 38–120 2·43–2·97	4.5-6.5 1.2-6.4 2.9-5.6 2.9-5.6
Sharma and Singh (2007)	$\begin{aligned} UCS &= 0.0642 \cdot V_{\rm p} - 117.99 \\ Id_2 &= 0.069 \cdot V_{\rm p} + 78.577 \end{aligned}$	$\begin{array}{c} 0.90 \\ 0.88 \end{array}$	7 types of rocks	10–1970 –	2-3.2.0
Kahraman and Yeken (2008)	$\rho = 0.213 \cdot V_{\rm p} + 1.256$	0.82	Carbonate rocks	2.0-2.6	3.6-6.1
This study	$UCS = 0.258 \cdot V_{p}^{3.543}$ $UCS = 49.4 \cdot V_{p} - 167$ $\rho = 0.19 \cdot V_{p} + 1.61$ $Id_{2} = 0.71 \cdot V_{p} + 95.7$	0.92 0.89 0.58 0.69	9 types of rock	20–125 – 2·15–2·85 –	1·89–6·1 - 1·8–6·1 -

**Table 1.** Relationship between  $V_p$  with both  $\rho$  and UCS.

\*UCS = uniaxial compressive strength (MPa),  $V_p$  = P-wave velocity (km/s),  $\rho$  = density (g/cm<sup>3</sup>)

Table 2. Geotechnical properties of studied rock units based on averaged properties and standard deviation.

Sampling locations	Rock type	$V_{\rm p}$ $\pm$ SD (km/s)	Hr $x \pm SD$	UCS $\overline{x} \pm SD$ (MPa)	$E \\ \overline{x} \pm SD \\ (GPa)$	$n' \\ \overline{x} \pm SD \\ (\%)$	$w \\ x \pm SD \\ (\%)$	$\rho_{\rm dry} \\ \overline{x} \pm {\rm SD} \\ ({\rm g/cm}^3)$	$\rho_{\rm sat} \\ \overline{x} \pm {\rm SD} \\ ({\rm g/cm}^3)$	$Id_2 \\ \overline{x} \pm SD \\ (\%)$
Denizli/Kocabas	Shrub travertine	$\begin{array}{c} 4 \cdot 8 \pm \\ 0 \cdot 12 \end{array}$	45 ± 4·5	61 ± 20·6	43 ± 6·9	$\begin{array}{c} 1.35 \pm \\ 0.46 \end{array}$	$0.55 \pm 0.19$	2·474 ± 0·25	2·488 ± 0·22	98.91 ± 0.10
Denizli/Kocabas	Noche travertine	$\begin{array}{c} 5 \cdot 0 \pm \\ 0 \cdot 08 \end{array}$	47 ± 3·1	64 ± 10·9	44 ± 3·3	$\begin{array}{c} 1\cdot 59 \pm \\ 0\cdot 89 \end{array}$	$\begin{array}{c} 0.66 \pm \\ 0.38 \end{array}$	$\begin{array}{c} 2 \cdot 419 \pm \\ 0 \cdot 48 \end{array}$	2·435 ± 0·42	98·55 ± 0·14
Denizli/Kaklık	Reed travertine	$\begin{array}{c} 4 \cdot 5 \pm \\ 0 \cdot 11 \end{array}$	39 ± 4·7	41 ± 16·6	35 ± 5·8	$\begin{array}{c} 1.89 \pm \\ 0.50 \end{array}$	$\begin{array}{c} 0.80 \pm \\ 0.22 \end{array}$	$\begin{array}{c} 2 \cdot 362 \pm \\ 0 \cdot 56 \end{array}$	2·381 ± 0·54	98·87 ± 0·12
Denizli/Honaz	Onyx travertine	$\begin{array}{c} 4 \cdot 7 \pm \\ 0 \cdot 19 \end{array}$	51 ± 3·3	58 ± 15	44 ± 5·2	$\begin{array}{c} 2 \cdot 05 \pm \\ 0 \cdot 88 \end{array}$	$\begin{array}{c} 0.76 \pm \\ 0.34 \end{array}$	$2.715 \pm 0.46$	$\begin{array}{c} 2\cdot735 \pm \\ 0\cdot38 \end{array}$	99·24 ± 0·07
Antalya/Korkuteli	Beige lime	$\begin{array}{c} 5 \cdot 0 \pm \\ 0 \cdot 17 \end{array}$	54 ± 1·9	$\begin{array}{c} 82 \pm \\ 28 \cdot 3 \end{array}$	$\begin{array}{c} 46 \pm \\ 4 \cdot 0 \end{array}$	$\begin{array}{c} 0.16 \pm \\ 0.10 \end{array}$	$\begin{array}{c} 0.06 \pm \\ 0.04 \end{array}$	$\begin{array}{c} 2{\cdot}682 \pm \\ 0{\cdot}08 \end{array}$	$\begin{array}{c} 2{\cdot}683 \pm \\ 0{\cdot}08 \end{array}$	99·43 ± 0·04
Denizli/Bozkurt	Dolomitic lime	$\begin{array}{c} 4 \cdot 9 \pm \\ 0 \cdot 29 \end{array}$	53 ± 2·4	92 ± 33·3	52 ± 11·8	$\begin{array}{c} 0.60 \pm \\ 0.27 \end{array}$	$\begin{array}{c} 0.22 \pm \\ 0.10 \end{array}$	2.778 ± 0.34	2·784 ± 0·32	99.65 ± 0.06
Antalya/Elmali	Soft lime	$3 \cdot 8 \pm 0 \cdot 41$	41 ± 3·6	32 ± 3·7	22 ± 4·7	$\begin{array}{c} 9.70 \pm \\ 2.20 \end{array}$	4·24 ± 1·14	$\begin{array}{c} 2 \cdot 311 \pm \\ 0 \cdot 98 \end{array}$	$\begin{array}{c} 2{\cdot}408 \pm \\ 0{\cdot}78 \end{array}$	98·49 ± 0·25
Denizli/Bekilli	Biotite schist	$5.1 \pm 0.44$	58 ± 4·4	98 ± 7·1	51 ± 7·9	$0.74 \pm 0.11$	$\begin{array}{c} 0.29 \pm \\ 0.04 \end{array}$	2.547 ± 0.42	$\begin{array}{c} 2 \cdot 554 \pm \\ 0 \cdot 43 \end{array}$	n/a
Denizli/Baklan	Mica schist	$5.6 \pm 0.32$	59 ± 1·1	114 ± 13·4	57 ± 6·8	$\begin{array}{c} 0.43 \pm \\ 0.53 \end{array}$	0·17 ± 0·21	2.638 ± 0.72	2.642 ± 0.67	n/a

 $\overline{x}$  = Average values and SD = Standard deviation.

ness plane, the results were excluded.  $V_{\rm p}$  is measured on samples by direct transmission using a Portable Ultrasonic Nondestructive Digital Indicating Tester (PUNDIT) that measures the time of propagation of ultrasound pulses with a precision of 0.1 µs and its transducers were 42 mm in diameter with 54 kHz (figure 2).  $V_{\rm p}$  test was performed perpendicular to observed layers. The *P*-wave velocity of studied rocks ranges from 3.8–5.1 km/s and can be classified as shown in table 3.

#### 3. Mineralogical composition of rock materials

Block rocks including travertine, limestone and schist were collected from various rock quarries in southwestern Turkey (figure 3). Mineralogical and textural studies were conducted on prepared thin section samples using optical microscope in accordance with EN 12407 (2002) Standard (table 4). Travertine, from Quaternary to Neogene ages, is one of the most common carbonate

rocks in the area. Travertine precipitated at different depositional conditions shows variations of colour, appearance, bedding, porosity, texture and composition (Chafetz and Folk 1984; Yagiz 2010). Travertine litho types in the basin mainly include shrub, onyx, reed and noche. Shrub type travertine represented by small bush like growths in the field is a common deposit on horizontal and sub horizontal surface in the basin (figure 3a). Noche, a commercial name for compact and dense reed type travertine is dark brownish in colour, dense and low porous (figure 3b). Reed travertine is one of the prominent elements in the study area and rich in molds of reed and coarse grass as in figure 3c (Guo and Riding 1998). Onyx travertine is commonly formed as a result of rapid precipitation due to fast flowing water on gentle slope. Dense, crudely fibrous and light coloured one is composed of elongated calcite feathers and developed perpendicular to the depositional surface (figure 3d). Jurassic aged beige coloured crystalline limestone and Eocene aged



Figure 1. Sampling locations in the study area.



Figure 2. *P*-wave velocity apparatus utilized in this study.

white coloured, fine grained spray-calcite cemented limestone outcrops around the City of Antalya (figures 3e and f). Dolomitic limestone is dark coloured, medium to coarse grained and massive in Eocene age (figure 3g). Two types of schist outcropped around the Denizli basin is categorized according to their mineralogical properties (figures 3h and i).

#### 4. Statistical analysis and discussions

Using the linear and nonlinear regression techniques including simple or multiple analysis for predicting the unknown from known variables are commonly encountered in the literature (O'Rourke 1989; Cargill and Shakoor 1990; Kahraman 2001; Sharma and Singh 2007; Yagiz 2008, 2009a). In this study, to develop the sets of empirical equations between  $V_{\rm p}$  and other rock properties including uniaxial compressive strength, modulus of elasticity, Schmidt hardness, slake durability index, effective porosity, water absorption by weight, and both saturated and dry density of rocks, linear and nonlinear simple regression analysis were performed with 95% confidence limits. To investigate the reliability of the obtained relationships, t-test and factor of significance (P-value) test were conducted among the achieved equations using the SPSS version 15 (2007) statistical package. Set of equations developed between the  $V_{\rm p}$  and measured rock properties using regression analysis are given in table 5. The significance of r-value can be determined by various statistical tests such as t-test and P-value test that is also known as observed level of significance test ( $\alpha$ ). The t-test compares the computed values with tabulated values using null hypothesis (Levine et al 2001). According to the *t*-test, when computed *t*-value is greater than tabulated *t*-value, the null hypothesis is rejected and obtained correlation coefficient (r-value) is acceptable. Also, observed level of significance is often used in hypothesis test. In this case, as *p*-value is smaller than level of significance ( $\alpha = 0.05$ ), the null hypothesis is rejected. Therefore, it means that there is a relation between the correlated parameters and this shows that r-value is significant.

In this study, the result of the regression analysis indicates that  $V_p$  have reliable relationship with the E, UCS, Hr, n', and water absorption by weight (figures 4–8) in accordance with the result of statistical analysis. On the other hand, the relation between  $V_p$  and Id<sub>2</sub> is not strong

 Table 3.
 P-wave velocity classification (Anon 1979).

$V_{\rm p}({\rm km/s})$	Description
	Very low Low Moderate High Very high



**Figure 3.** Mineralogical and textural studies of rock using thin section analysis ( $\times$  10); (**a**) Shrub travertine (**b**) Noche travertine (**c**) Reed travertine (**d**) Onyx travertine (**e**) Beige limestone (**f**) Soft limestone (**g**) Dolomitic limestone (**h**) Biotite schist (**i**) Muscovite schist.

Table 4.	Mineralogical an	nd petrographical	composition of studied	l rocks together with lo	cation.
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Rock type/ lithology	Locations	Microscopic description	Grain size
Sedimentary/	Denizli/	Sparite micrite cemented, densely packed texture,	Fine to fine
shrub travertine	Kocabas	crystals 5–10 μm in size with no internal architecture	medium
Sedimentary/	Denizli/	Sparite calcite cemented. Crystal size 20 µm or more in dia;	Fine to
Noche travertine	Kocabas	compact texture with low pores and organic matter.	medium
Sedimentary/	Denizli/	Sparite calcite cemented, crystal size from 20–150 µm in diameter. Relatively more organic content and porous.	Fine to
reed travertine	Kaklik		medium
Sedimentary/ onyx travertine	Denizli/ Honaz	Micrite and sparite cemented matrix, dark coloured. Sparry crystals $10 \ \mu m$ wide and $100-200 \ \mu m$ in length with iron content, layered texture.	Fine with organic matter
Sedimentary/ beige limestone	Antalya/ Korkuteli	Sparite calcite cemented, micro crack observed with coarse spar-calcite and calcite fillings, no fossils.	Medium to coarse
Sedimentary/ dolomitic lime	Denizli/ Bozkurt	Sparite and micrite cemented texture with healed joints filled with secondary calcite fillings.	Medium to coarse
Sedimentary/	Antalya/	Sparite calcite cemented, light cream coloured,	Fine grain
white limestone	Elmali	some micro fossils with no cracks.	
Meta-sedimentary/ quartz-biotite-schist	Denizli/ Bekilli	Crystal size 0.1–0.2 mm, elongated calcite and quartz crystals along the schistose, quartz biotite, mica schist with opaque and schistose texture.	Fine to medium fine
Meta-sedimentary/	Denizli/	Crystal size 0·1–0·2 mm, elongated calcite and quartz crystals along the schistose, quartz muscovite, schist with opaque and schistose texture.	Fine to
quartz-mica-schist	Baklan		medium fine

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Rock properties	Equations	<i>r</i> -value	<i>t</i> -value	<i>t</i> -table	$P$ -value < $\alpha = 0.05$
E (GPa)	$E = 20 \cdot 1 \cdot V_{\rm p} - 53$	0.95	8.22	$\pm 2.31$	0.000
UCS (MPa)	$UCS = 49.4 \cdot V_p - 167$	0.89	5.22	$\pm 2.31$	0.001
n' (%)	$n' = -5 \cdot 19 \cdot V_{\rm p} + 27 \cdot 1$	0.86	-4.53	$\pm 2.31$	0.003
w (%)	$w = -2 \cdot 23 \cdot V_{\rm p} + 11 \cdot 6$	0.85	-4.30	$\pm 2.31$	0.004
Hr	$Hr = 11.68 \cdot V_p - 6.64$	0.80	3.55	$\pm 2.31$	0.009
Id <sub>2</sub>	$Id_2 = 0.71 \cdot V_p + 95.7$	0.69	2.12	$\pm 2.31$	0.088
$\rho_{\rm dry}  ({\rm g/cm}^3)$	$\rho_{\rm dry} = 0.19 \cdot \dot{V}_{\rm p} + 1.61$	0.58	1.85	$\pm 2.31$	0.107
$\rho_{\rm sat} ({\rm g/cm^3})$	$\rho_{\rm sat} = 0.14 \cdot V_{\rm p} + 1.88$	0.46	1.40	$\pm 2.31$	0.206

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**Table 5.** Empirical equations between  $V_{\rm p}$  and the measured rock properties.



**Figure 4.** Relationship between the UCS and  $V_{\rm p}$ .







Figure 5. Relationship between the E and  $V_{p}$ .

enough to rely on (r = 0.69) as given in figure 9. Further, the  $V_p$  provides lower relationship with dry and saturated

**Figure 7.** Relationship between the n' and  $V_{p}$ .

density of rock with r = 0.58 and 0.46 respectively (figure 10). To evaluate the validity of the generated equations between the rock properties and *P*-wave velo-

city, obtained empirical equations are compared with those equations available in the literature. The relationship between  $V_{\rm p}$  and rock properties including effective porosity, water absorption by weight and slake durability index



**Figure 9.** Relationship between  $Id_2$  and  $V_p$  together with previous study from literature.



**Figure 10.** Relationship between both the  $\rho_{dry}$  and  $\rho_{sat}$  with  $V_p$ .

are compared with the previous studies in figures 7–9. Further, produced empirical relationship between  $V_{\rm p}$  and both UCS and dry density of rocks are also associated with previous researches in figures 11 and 12. As seen from the figures 7–12, the obtained equations and coefficient of correlations are various ranging from 0.46 to 0.95.

## 5. Conclusions

Using the standard testing methods, nine different rock types were tested and the results examined to generate



Figure 11. Comparison of the obtained results with previous researches;  $\rho_{dry}$  vs  $V_{p}$ .



**Figure 12.** Comparison of the obtained results with previous researches; UCS vs  $V_{p}$ .

empirical relationship between  $V_p$  and rock properties including uniaxial compressive strength, modulus of elasticity, Schmidt hardness, slake durability index, effective porosity, water absorption by weight and both dry and saturated density of rocks. Further, petrographical and mineralogical studies were conducted using thin section analysis. It is evident that rock texture and type have great affect on their geotechnical properties. The result shows that the UCS, E, Hr, n', w and Id<sub>2</sub> of rocks can be estimated by conducting  $V_{\rm p}$  test that is non-destructive, simple, faster and a relatively economic method for rock characterization. The best relationships obtained between the  $V_{\rm p}$  and both UCS and E were with correlation coefficient of 0.92 and 0.95, respectively. The relationship between the  $V_{\rm p}$  and measured rock properties are acceptable according to the statistical analysis including *t*-test, *p*-value test and coefficient of correlations except that obtained between the  $V_{\rm p}$  and both dry and saturated unit weight of rock (r = 0.58 and 0.46, respectively). The investigated rock properties excluding dry and saturated density of rock can be estimated as function of  $V_{\rm p}$  using derived equations; however, those equations should be used with care for only similar rocks.

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